

Numerical Modeling of Stratified Shallow Flows: Applications to Aquatic Ecosystems

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Outline

- 1 Motivations
- 2 A variable-density multilayer Saint-Venant model
- 3 A Reservoir model based on Navier–Stokes: OPHÉLIE^{©EDF}
- 4 Numerical simulations
- 5 Perspectives

Modeling of complex free surface environmental flows

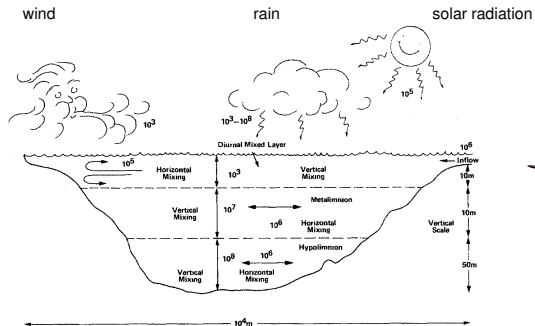
Aquatic ecosystems: lakes, estuarine waters, lagoons,...

Related applications: water quality, sustainability of ecosystems.

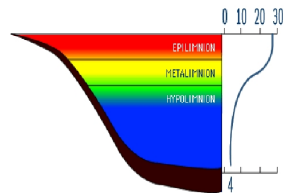


- Important vertical **stratification** (temperature, salinity);
- Column structure result of complex 3D forcing mechanisms:
 - Natural: rivers inflows, wind, solar radiation,...
 - Man-made: inflow/outflow in hydraulic reservoirs, pollutant spills,...
- Bio-dynamics primarily controlled by thermo-chemical status.

Typical lake stratification



Lake stratification with estimates of energy transfer time scales [s]
[J. Imberger, *Hydrobiologia* 125, 7–29, 1985]



Temperature vertical profile [°C]
(<http://waterontheweb.org>)

A Multilayer Saint-Venant Model with Variable Density

- Saint-Venant models

Main assumption: $\varepsilon = H_0/L_0 \ll 1$ (shallow flow)

- Classical Saint-Venant equations

- Uniform density and vertically uniform velocity;
- Hydrostatic pressure (vertical velocity = 0).

Multilayer Saint-Venant models

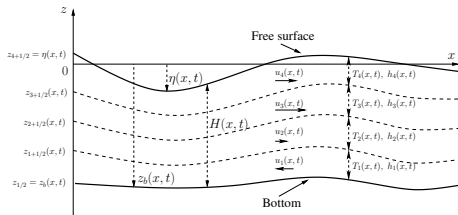
- Allow non-uniform velocity along the vertical.
 - Immiscible layers [Ovsyannikov 79, Vreugdenhil 79, Parés et al. 00,01,03 ...]
 - Multilayer model with mass exchange
[Audusse–Bristeau–Perthame–Sainte-Marie, 2009]
- Allows fluid circulation between layers.

Goal: Introduce variable density description via the multilayer strategy.

Variable-Density Multilayer Saint-Venant System

Hydrostatic Euler equations with $\rho = \rho(T)$, $T = \text{active tracer}$

$$\begin{aligned} \frac{\partial \rho}{\partial t} + \frac{\partial(\rho u)}{\partial x} + \frac{\partial(\rho w)}{\partial z} &= 0, \\ \frac{\partial(\rho u)}{\partial t} + \frac{\partial(\rho u^2)}{\partial x} + \frac{\partial(\rho u w)}{\partial z} + \frac{\partial p}{\partial x} &= 0, \\ \frac{\partial p}{\partial z} &= -\rho g, \\ \frac{\partial(\rho T)}{\partial t} + \frac{\partial(\rho u T)}{\partial x} + \frac{\partial(\rho w T)}{\partial z} &= 0. \end{aligned} \quad + \text{B.C.}$$



$$h_\alpha = l_\alpha H, \quad \sum_{\alpha=1}^N l_\alpha = 1$$

$$u(x, z, t) \approx \sum_{\alpha=1}^N 1_{z \in h_\alpha}(z) u_\alpha(x, t)$$

$$T(x, z, t) \approx \sum_{\alpha=1}^N 1_{z \in h_\alpha}(z) T_\alpha(x, t)$$

Variable-Density Multilayer Saint-Venant System

Unknowns: $H, u_\alpha, T_\alpha, \alpha = 1, \dots, N$

$$\frac{\partial}{\partial t} \left(\sum_{\alpha=1}^N \rho_\alpha h_\alpha \right) + \frac{\partial}{\partial x} \left(\sum_{\alpha=1}^N \rho_\alpha h_\alpha u_\alpha \right) = 0,$$

$$\begin{aligned} \frac{\partial}{\partial t} (\rho_\alpha h_\alpha u_\alpha) + \frac{\partial}{\partial x} (\rho_\alpha h_\alpha u_\alpha^2 + h_\alpha p_\alpha) \\ = u_{\alpha+1/2} G_{\alpha+1/2} - u_{\alpha-1/2} G_{\alpha-1/2} + \frac{\partial z_{\alpha+1/2}}{\partial x} p_{\alpha+1/2} - \frac{\partial z_{\alpha-1/2}}{\partial x} p_{\alpha-1/2}, \end{aligned}$$

$$\begin{aligned} \frac{\partial}{\partial t} (\rho_\alpha h_\alpha T_\alpha) + \frac{\partial}{\partial x} (\rho_\alpha h_\alpha u_\alpha T_\alpha) \\ = T_{\alpha+1/2} G_{\alpha+1/2} - T_{\alpha-1/2} G_{\alpha-1/2}, \quad \alpha = 1, \dots, N. \end{aligned}$$

$$h_\alpha = l_\alpha H, \quad z_{\alpha+1/2}(x, t) = z_b(x) + \sum_{j=1}^{\alpha} h_j(x, t),$$

$$p_\alpha = \frac{g}{2} \rho_\alpha h_\alpha + g \sum_{j=\alpha+1}^N \rho_j h_j, \quad p_{\alpha+1/2} = g \sum_{j=\alpha+1}^N \rho_j h_j,$$

$$G_{\alpha+1/2} = \frac{\partial}{\partial t} \left(\sum_{j=1}^{\alpha} \rho_j h_j \right) + \frac{\partial}{\partial x} \left(\sum_{j=1}^{\alpha} \rho_j h_j u_j \right), \quad \rho_\alpha = \rho(T_\alpha).$$

- **Hyperbolicity ?** Typically hyperbolic in practice.

Numerical Method

System of $2N+1$ equations in $X = (H, \{T_\alpha, q_\alpha\}_{1 \leq \alpha \leq N})$, $q_\alpha = \rho_\alpha h_\alpha u_\alpha$:

$$\frac{\partial}{\partial t} W(X) + \frac{\partial}{\partial x} f(X) = S_G(X, \partial_{t,x} X) + S_p(X) .$$

$\uparrow_{G_{\alpha \pm 1/2}} \qquad \qquad \qquad \uparrow_{p_{\alpha \pm 1/2}}$

- Conservative portion and S_G : **Finite Volume Kinetic Scheme** [Perthame–Simeoni, 2001]
 - Positivity preserving;
 - Avoids eigenvalues computation.
- Source S_p : **Hydrostatic Reconstruction Method** [Audusse et al., 2004]
 - **Well-balancing** of equilibrium states. Particular case at rest:

$$u_\alpha = 0, \quad H + z_b = \text{const.}, \quad \rho_\alpha = \text{const.}, \quad \alpha = 1, \dots, N.$$

Additional implemented features:

- 2nd order in space (MUSCL) and time (Heun);
- Vertical viscosity for velocity and temperature (implicit method).

OPHÉLIE

Hydrodynamic reservoir model based on Navier–Stokes.

[M.-J. Salençon, J.Y. Simonot, 1989] © EDF

2D laterally averaged hydrostatic incompressible Navier–Stokes eqs.

- Rigid lid hypothesis: $\frac{\partial z_s}{\partial x} = \frac{\partial z_s}{\partial y} = 0$.

$$\frac{\partial(Bu)}{\partial x} + \frac{\partial(Bw)}{\partial z} = 0, \quad B(x, z) = \text{lake width},$$

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + w \frac{\partial u}{\partial z} = \frac{1}{B} \frac{\partial}{\partial x} (\nu_x B \frac{\partial u}{\partial x}) + \frac{1}{B} \frac{\partial}{\partial z} (\nu_z B \frac{\partial u}{\partial z}) - \frac{\partial p}{\partial x} - g \frac{u|u|}{BC^2},$$

$$\frac{\partial T}{\partial t} + u \frac{\partial T}{\partial x} + w \frac{\partial T}{\partial z} = \frac{1}{B} \frac{\partial}{\partial x} (D_x B \frac{\partial T}{\partial x}) + \frac{1}{B} \frac{\partial}{\partial z} (D_z B \frac{\partial T}{\partial z}) - \frac{1}{\rho_0 C_p} \frac{\partial Q}{\partial z},$$

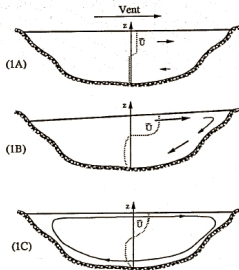
$$p = p_s + g\rho_0(z_s - z) - \kappa g\rho_0 \int_z^{z_s} (T - T_0)^2 dz, \quad \rho(T) = \rho_0(1 - \kappa(T - T_0))^2.$$

- ★ Includes surface thermodynamics.
- ★ Rigid lid vs. free-surface: much less restrictive CFL constraint.
(BUT limited range of applications.)

Model developed for specific application: Lake of Sainte-Croix.

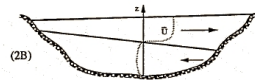
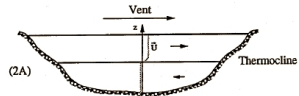
Lake response to wind stress

wind
→

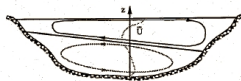


LAC HOMOGENE

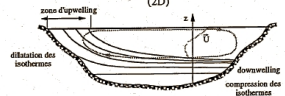
Homogeneous lake



Vent faible - stratification forte



Vent fort - stratification faible



LAC STRATIFIE

Stratified lake

M.-J. Salençon and J.-M. Thébault, *Modélisation d'écosystème lacustre*, 1997.

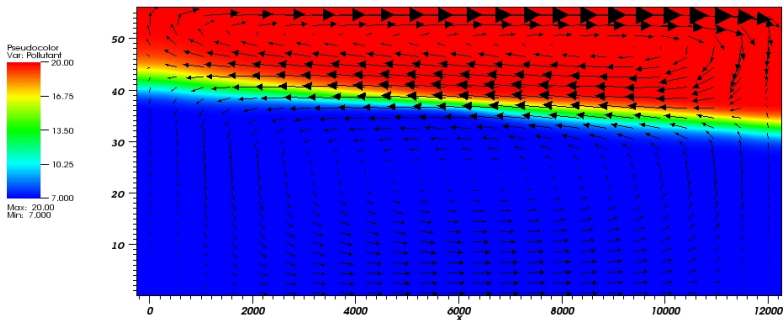
(After J. Imberger, 1979, 1987).

Test 1

$V_W = 10$ m/s, $\Delta H_{\text{meta}} = 2.5$ m.

Uplifting of the thermocline until steady position.

$t = 82827.9$

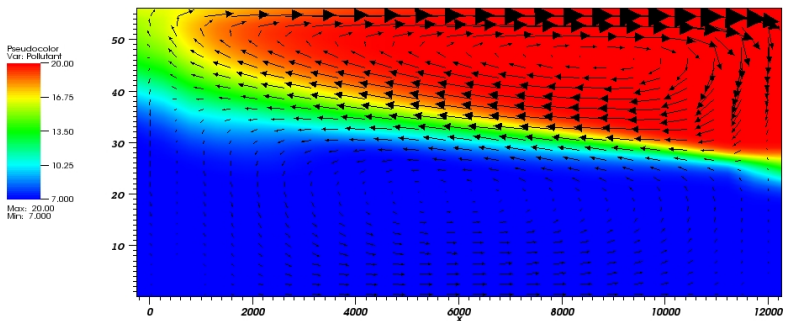


Test 2: Upwelling

$$V_W = 15 \text{ m/s}, \Delta H_{\text{meta}} = 20 \text{ m.}$$

Upwelling of the middle layer (metalimnion).

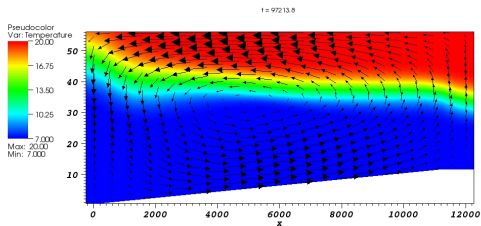
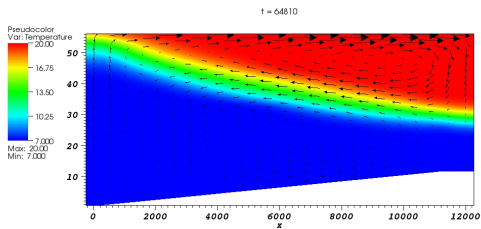
$t = 172816$



Test 3: Variable topography

$V_W = 15$ m/s, $\Delta H_{\text{meta}} = 20$ m.

Simulation with wind stopped after 18 hours.

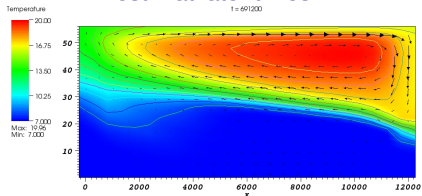


Limitations of the hydrostatic assumption

- Model lacks description of **unstable stratification**.

Static equilibrium stable if $\frac{\partial \rho}{\partial z} \leq 0$.

Test 2 at later times:



Whereas physically:
cold water front propagation
with vertical isotherms.
(cf. [results OPHÉLIE](#))

Non-hydrostatic terms need to be included in the model:

$$\frac{\partial}{\partial t}(\rho w) + \frac{\partial}{\partial x}(\rho u w) + \frac{\partial}{\partial z}(\rho w^2) + \frac{\partial p}{\partial z} = -\rho g.$$

- Preliminary results with modeling of $\frac{\partial}{\partial z}(\rho w^2)$ (Exp. Rayleigh).

OPHÉLIE: Stabilization of temperature profile by **mixing**.

Current and future work

Current studies

- Variable-density multilayer Saint-Venant model
 - Non-hydrostatic terms.
 - Surface thermodynamics.

Prospected work

- Comparison multilayer S.-V. / OPHÉLIE (time scale \sim days).
- Coupling hydrodynamics with **biological processes**.
- OPHÉLIE application to realistic environment (Lac de Grangent).